

# **A Modeling Study of Interannual to Decadal Variability in Equatorial Pacific Biogeochemistry and Ecosystems**

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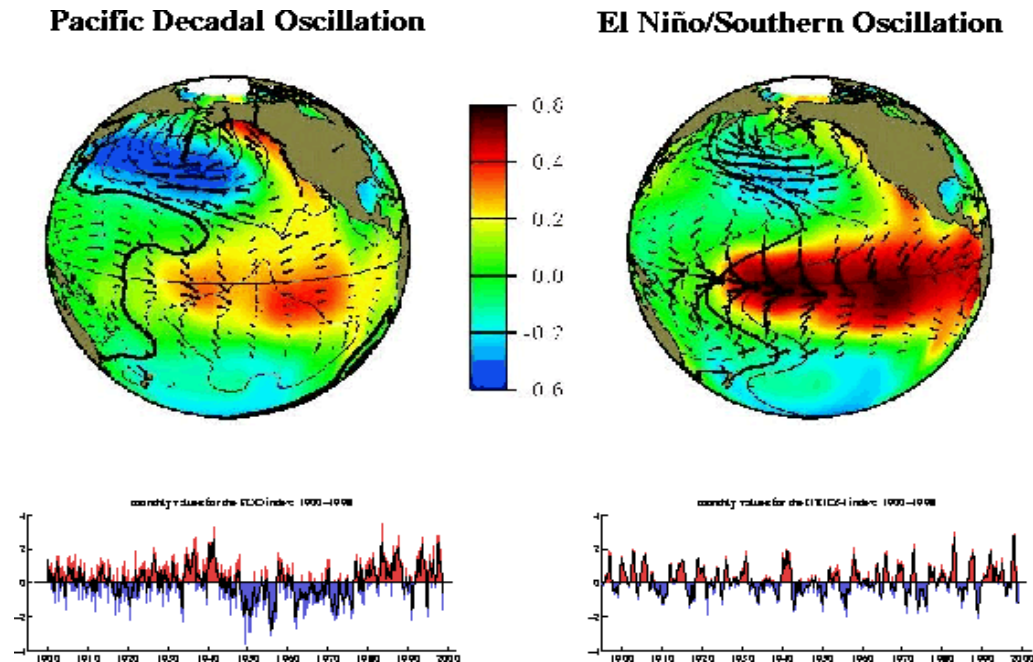
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Christophe Mekes (LOCEAN, Paris)

Thomas Gorgues (LOCEAN, Paris)

# The Pacific Decadal Oscillation (PDO)

- Timeseries: construction by low-pass filtering. The PDO is often interpreted as a “slow mode” moving through system



Mantua et al. [1997]

- McPhaden and Zhang (2002) : Changes in overturning of “subtropical cells” (STCs) which link equatorial upwelling to extratropical subduction regions. Equatorial upwelling reduced by about 25%

## **General Question:**

**For model simulations, how do Equatorial Pacific biogeochemistry and ecosystems respond to wide range of timescales of variability contained in NCEP reanalysis fluxes?**

## **Focus:**

**Decadal variability: Do the slowdown of the STCs and the resulting reduction in upwelling impact on equatorial biogeochemistry?**

**Is there a biogeochemical connection between extratropical and tropical basins through the ocean?**

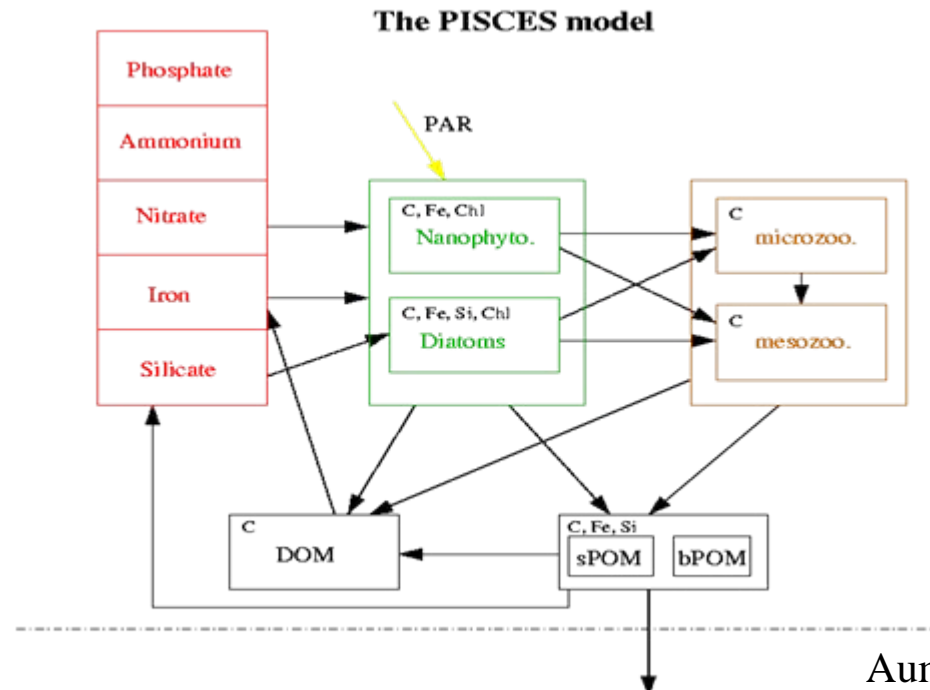
*Rather than work with anomalies of tracer fields, consider full monthly mean tracer fields in order to identify relationship between different timescales*

# MODEL CONFIGURATION

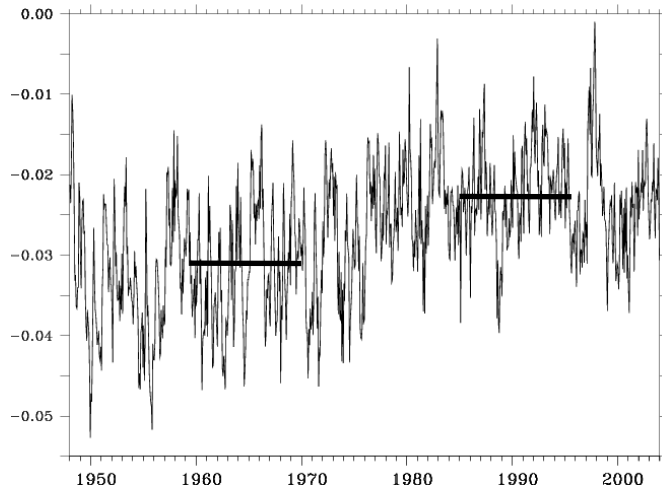
## ■ Circulation

Forced experiments with ORCA2 configuration of OPA 8.2 model (Madec et al. 1998); model has 30 vertical layers, 2° resolution in extratropics, and 0.5° meridional resolution in the equatorial regions. Surface buoyancy fluxes are calculated using bulk formulas. Surface forcing from NCEP over 1948-2003.

## ■ Biogeochemistry



# Decadal changes in NCEP wind stresses



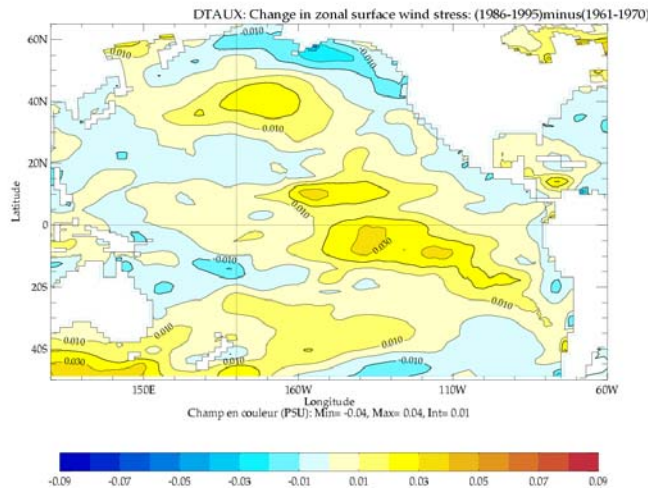
NCEP  $\tau_x$  (zonal wind stress component) averaged over 4°N-4°S, 120°E-270°E,. This quantity is 34% larger in NCEP over the period 1961-1970 than 1986-1995.

NCEP: despite some problems, best climate reconstruction that we have. How realistic is the 1976/77 “shift”, and to what extent is it an artifact of assimilation of satellite data?

Sigfried Schubert (NASA): shift not due to data assimilation

Difference in surface wind stress pattern:  
Mean(1986-1995)-Mean(1961-1970)

Positive values indicate that wind stresses were weaker for the later period.

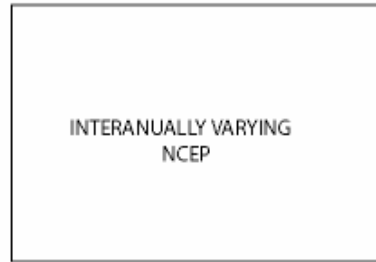


## PHYSICAL FORCING FOR 3 EXPERIMENTS

### CONTROL EXPERIMENT: PISCCTL

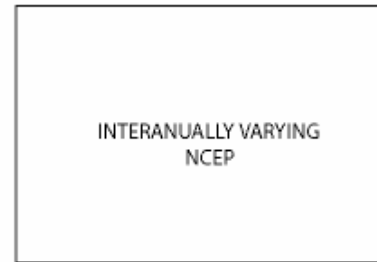
WINDSTRESS FORCING:

Y



BUOYANCY FORCING:

Y

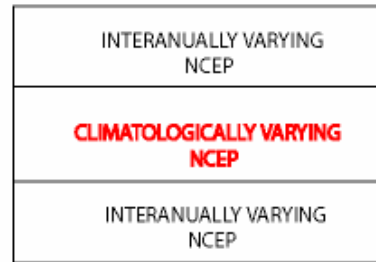


### CLIM\_EQTAU EXPERIMENT

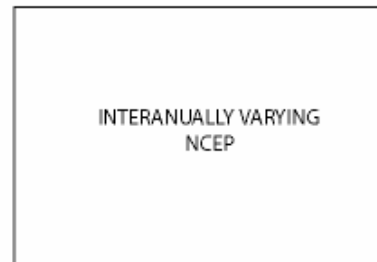
WINDSTRESS FORCING:

15°N

15°S



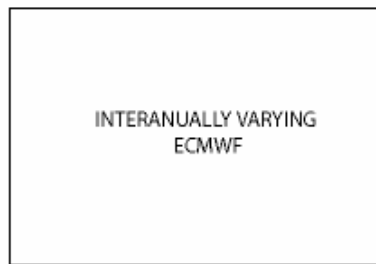
BUOYANCY FORCING:



### PISCERA EXPERIMENT

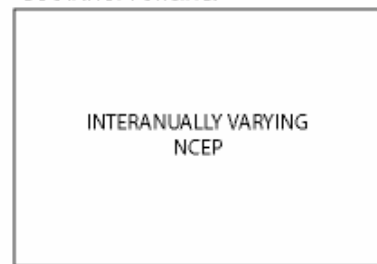
WINDSTRESS FORCING:

Y



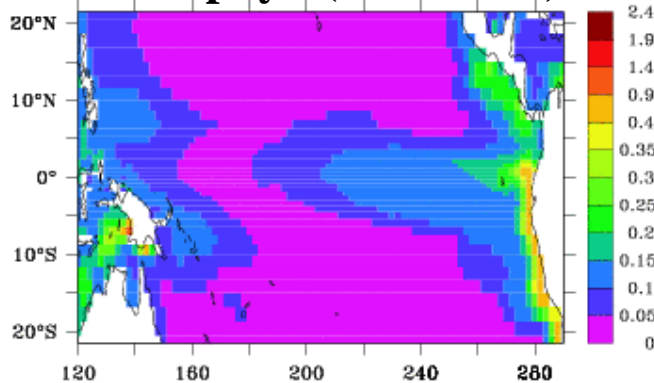
BUOYANCY FORCING:

Y



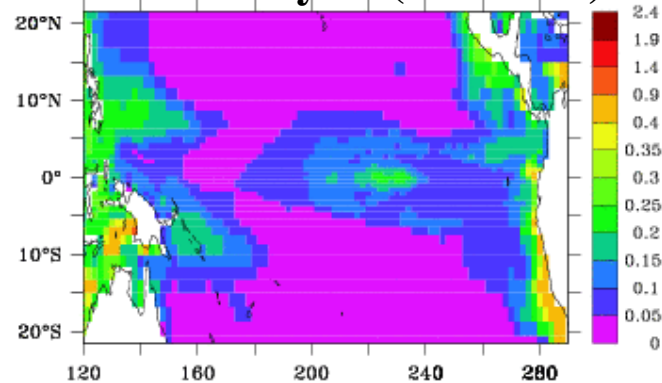
# MODEL/DATA COMPARISON FOR SURFACE CHLOROPHYLL:

**Mean ORCA2-PISCES  
chlorophyll (1998-2002)**

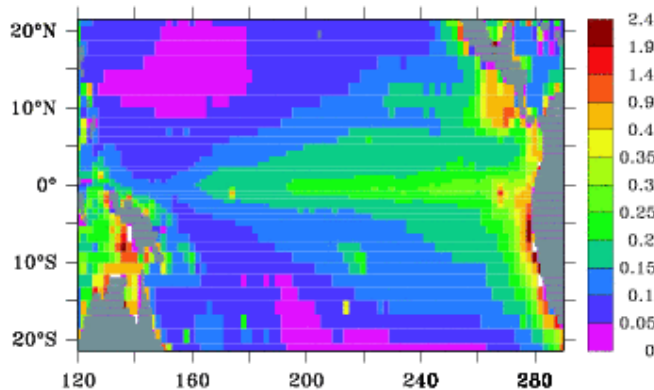


**ORCA2-PISCES**

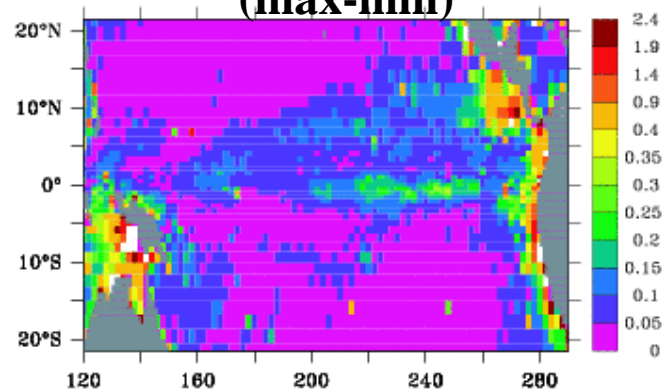
**seasonal cycle (max-min)**



**Mean SEAWIFS chlorophyll**



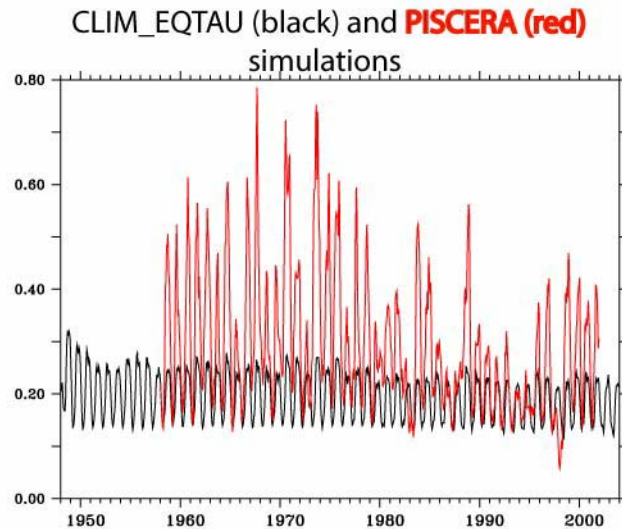
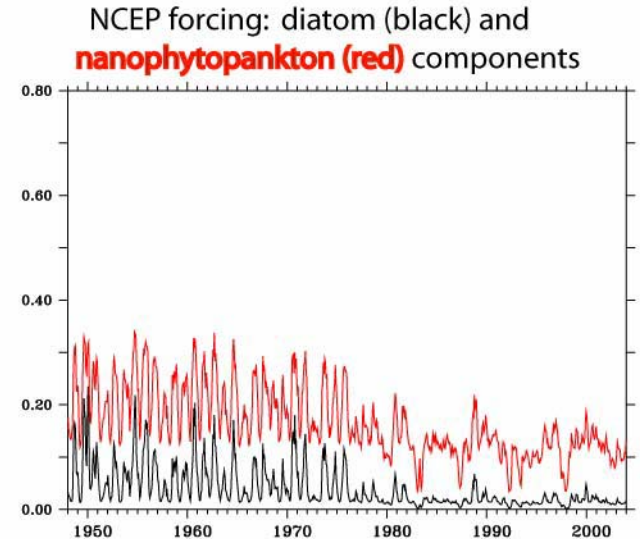
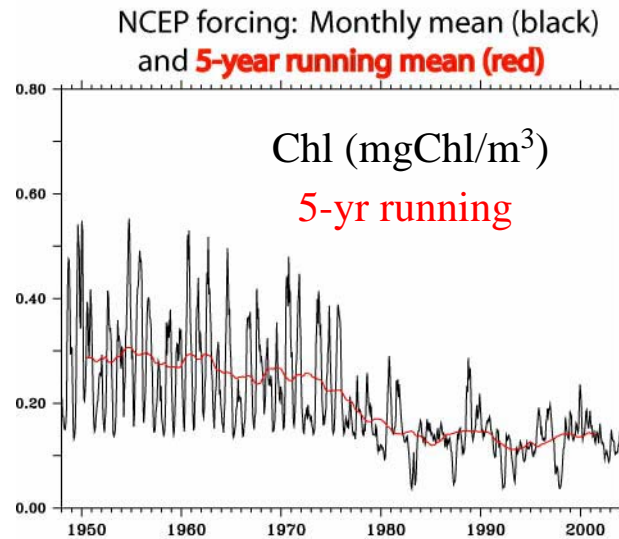
**SEAWIFS seasonal cycle  
(max-min)**



(Units of  $\mu\text{mol/L}$ )

# Monthly Chlorophyll Concentration averaged over 90W-150W,2N-2S (1948-2003)

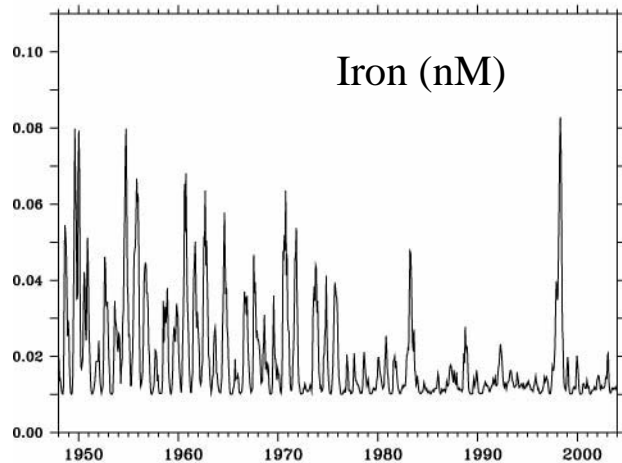
61-70 vs. 86-95  
Annual: 80% higher  
Seasonal: 120% higher



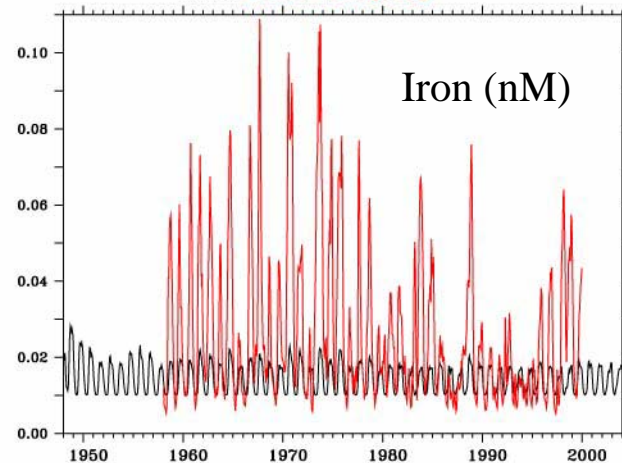


# Monthly Fe and NO<sub>3</sub> concentrations averaged over 90W-150W,2N-2S (1948-2003)

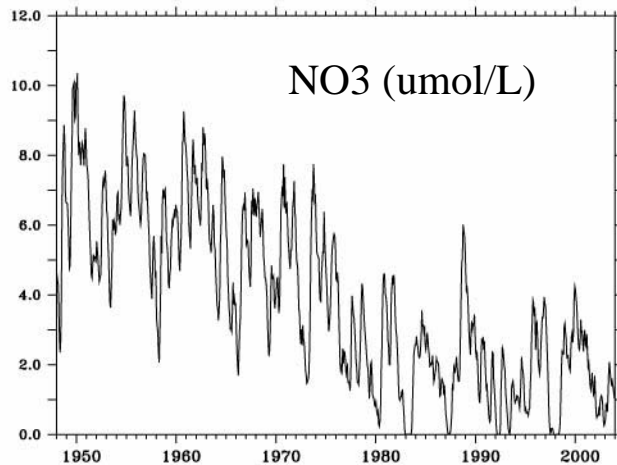
Fe concentration for NCEP forcing



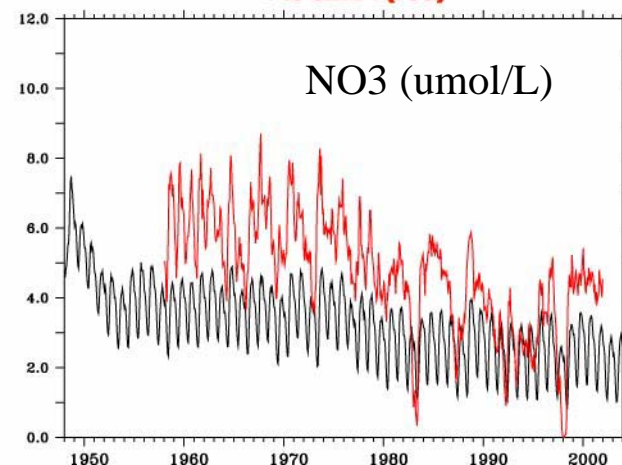
Fe concentration for CLIM\_EQTAU  
and **PISCERA (red)**



NO<sub>3</sub> concentration for NCEP forcing



NO<sub>3</sub> concentration for CLIM\_EQTAU  
and **PISCERA (red)**

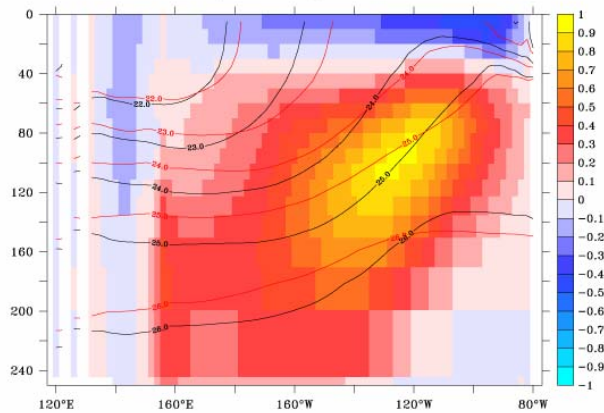


## Potential mechanisms

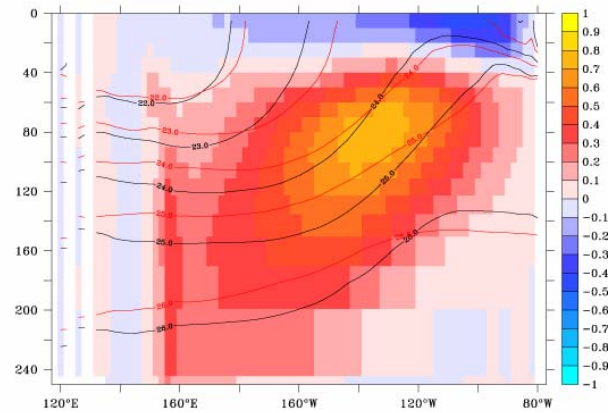
- **Change in upwelling rate (volume/time)**
- **Change in depth (or isopycnal horizon) from which upwelling occurs, with iron concentration remaining roughly constant on isopycnal surfaces**
- **Ferricline/pycnocline decoupling (change in iron concentration on isopycnal surfaces)**

### Decadal changes in zonal and vertical velocity along equator for Pacific (1961-1970 and 1986-1995)

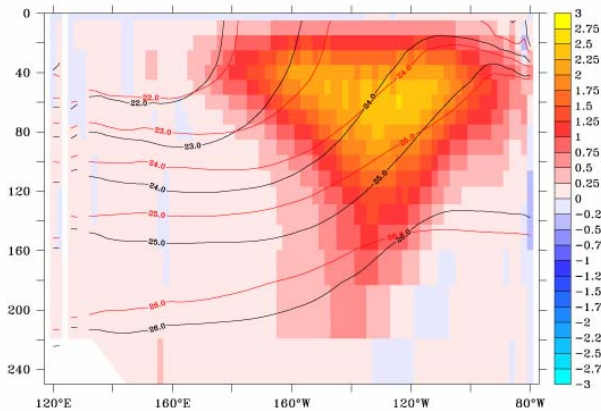
Zonal velocity: average over 1961-1970



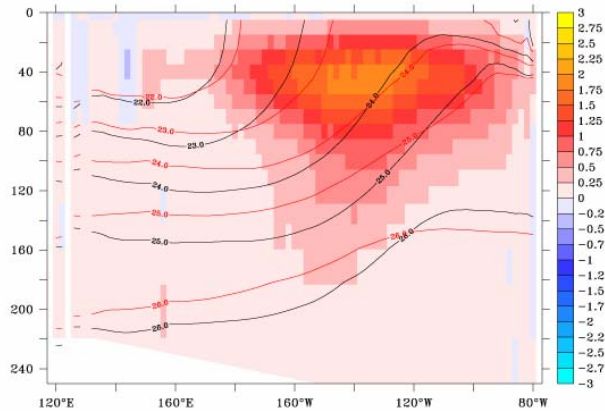
Zonal velocity: average over 1986-1995



Vertical velocity: average over 1961-1970



Vertical velocity: average over 1986-1995

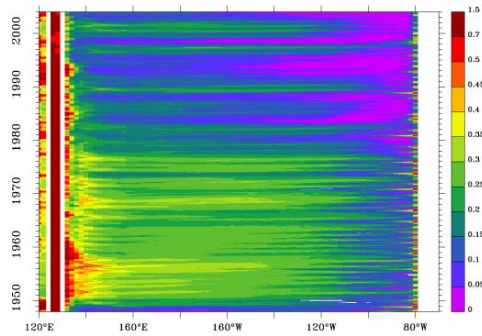


# Tracers on isopycnal surfaces along Equator (1948-2003)

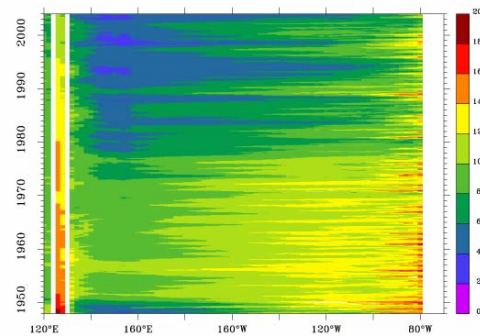
HOVMOLLER DIAGRAMS OF EVOLUTION OF TRACER  
CONCENTRATION ALONG EQUATOR OVER 1948-2003

PISCTL

Fe on  $\sigma_0=25.0$  for PISCINT (1948-2003)

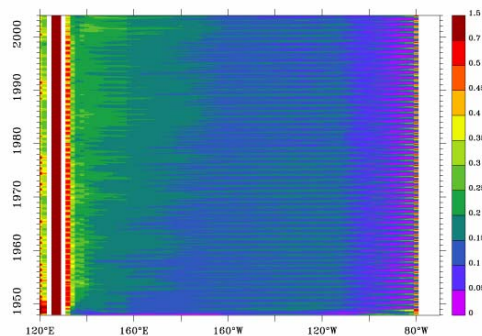


NO3 on  $\sigma_0=25.0$  for PISCINT (1948-2003)

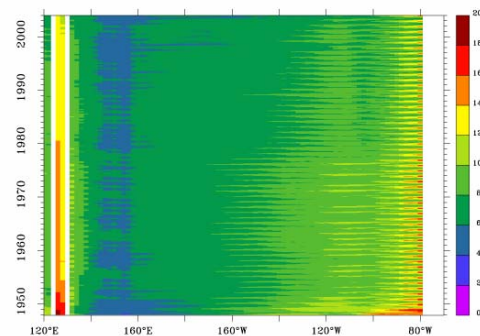


CLIM\_EQTAU

Fe on  $\sigma_0=25.0$  for CLIM\_EQTAU (1948-2003)

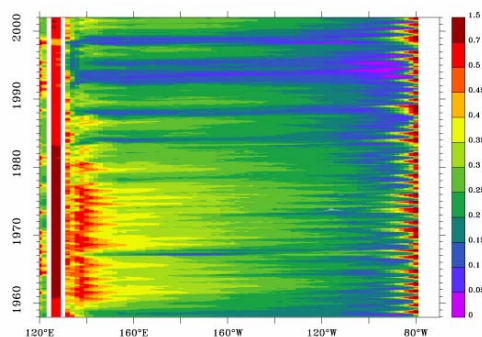


NO3 on  $\sigma_0=25.0$  for CLIM\_EQTAU (1948-2003)

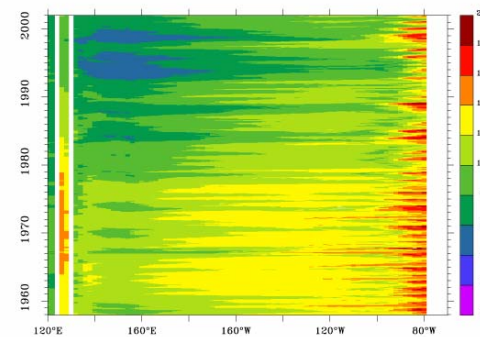


PISCERA

Fe on  $\sigma_0=25.0$  for PISCINT (1948-2003)



NO3 on  $\sigma_0=25.0$  for PISCINT (1948-2003)

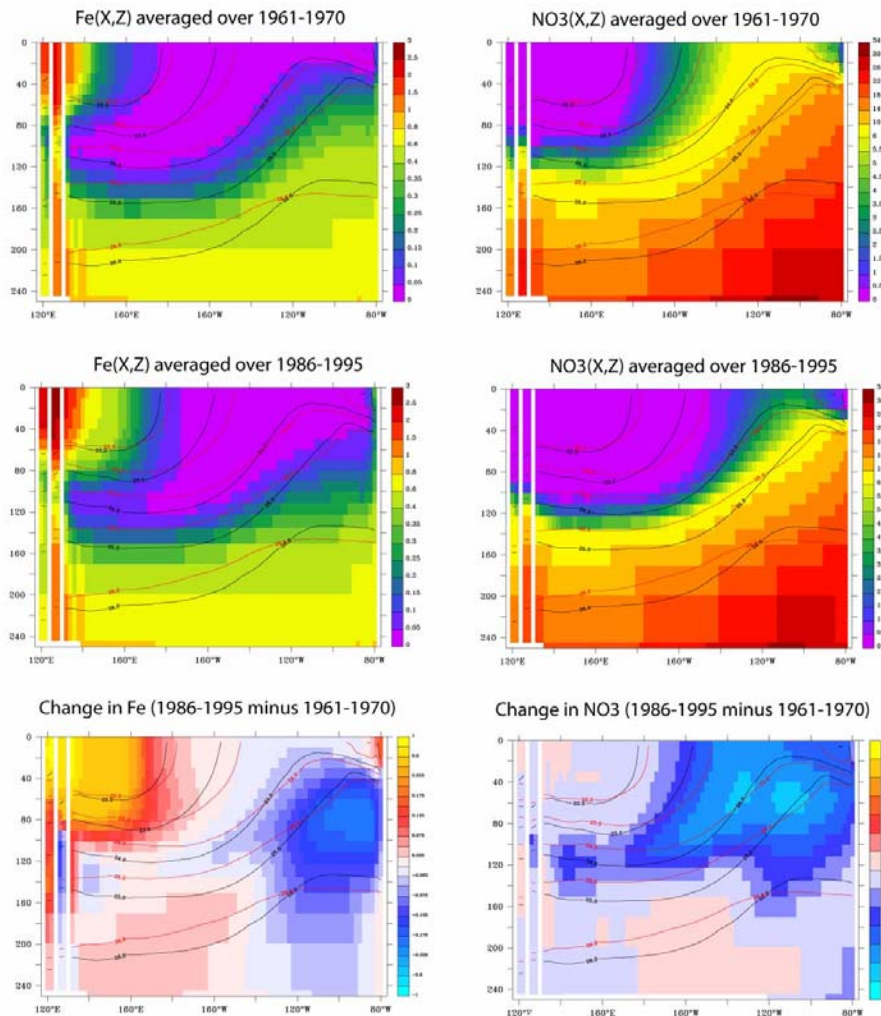


Fe (nmol/kg)

NO3 ( $\mu\text{mol/kg}$ )

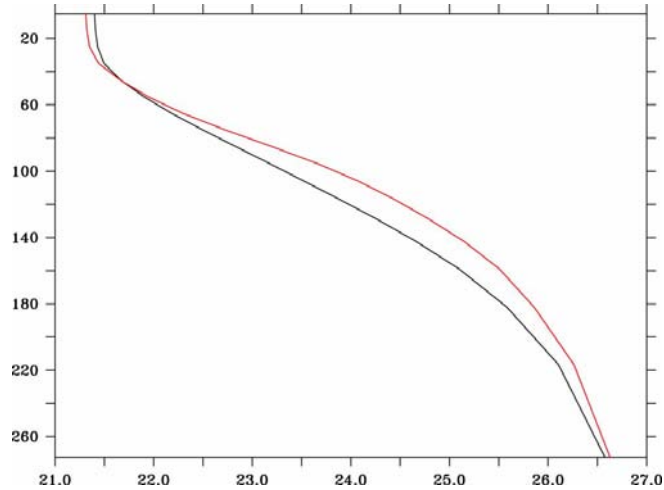


# Decadal change in tracer concentrations $\square$ along Equator (1961-1970 and 1986-1995)

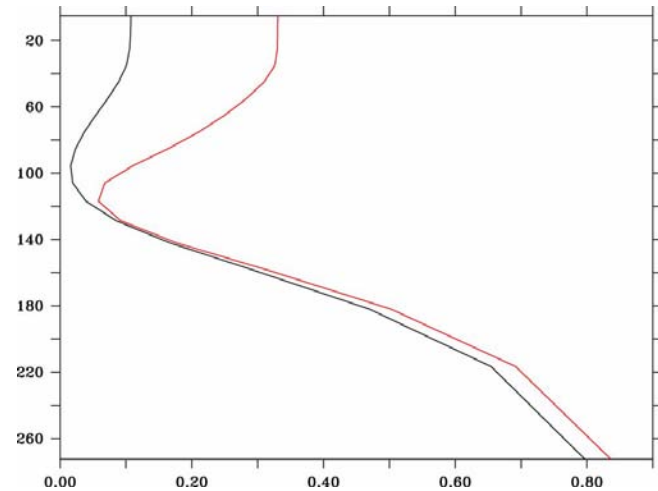


# VERTICAL PROFILES OF DENSITY AND DISSOLVED IRON CONCENTRATION UNDER THE WARM POOL (Average over 155E-165E, 1N-1S)

Potential density



Iron concentration



Decade 1961-1970

Decade 1986-1995

# Conclusions

- Large decadal changes in seasonality of surface iron and chlorophyll
- Surface Chlorophyll response strongly nonlinear (amplification here)
- Decoupling of ferricline and pycnocline on decadal timescales

## Is it realistic ?

- We don't know. Very few data before 1976 for Chl and NO<sub>3</sub>
- Strong blooms possible (La Nina 98). Decadal changes possible (Antoine et al., 2005, ...)
- Main conclusions:

Equatorial Pacific may not be stable (low Chl, low Seasonality)

Biogeochemical responses are non-linear

Important of not overfiltering the data (seasonality)